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There are two types of the magnetic field features are the magnetic field the surrounding area. Some features are always present within these areas but they always have a lower incidence than the surrounding area. These low intensity areas are important and it is quite likely that they are due to some subsurface feature. The type of feature will be discussed in the next section of this report.

STRUCTURE

Petroleum and Natural Gas Permit No.5059 is located on the interior plain of the Northwest Territories about 50 miles from the west of the edge of the Pre-Cambrian Shield. The strike of the sedimentary rocks is about north 30 degrees west and the units dip to the southwest at a few tens of feet per mile.

Structural features which could be present and which could cause the low incidence anomalies mentioned in this report are discussed in order of probability.

(1) PRE-CAMBRIAN TOPOGRAPHY

Basement topography under Permit No.5059 is thought to be much the same as it is today along the southwest edge of the Shield. Low rounded hills separated by gentle to abrupt valleys are seen on the Shield and these

features are undoubtedly present under the subject Permit. The effect of this Basement relief on the overlying sedimentary rocks is often great. The Granite Wash sand is usually present in the topographic "lows" on the Basement but absent on the "highs". The Granite Wash is an excellent potential reservoir.

Further effects of Basement topography on beds higher than the Granite Wash is gentle folding present over Basement hills. These folds are anticlines in every sense and could form traps for oil or gas.

Many small faults have been reported by A.W. Norris (1965) in the Basement and immediately overlying rocks and these features could cause closure within the sedimentary units.

2. DEVONIAN REEFS

Devonian reefs strongly affect the fracture pattern and control the occurrence of gas and oil in the overlying beds. Devonian reefs are present west of this Permit and others could well be present under the subject area.

3. TECTONIC FOLDING & FAULTING

The presence of tectonic folds is very unlikely, but some normal faulting is probably present.

4. TOPOGRAPHY RELIEF ON AN INTRA-SEDIMENTARY UNCONFORMITY

Unconformity, is a possible source of fracture intensity anomalies, but within the Permit area it is unlikely that the relief on any unconformities within the sedimentary section is great enough to affect the fracture pattern.

Reference to the Total Fracture Pattern Map which accompanies this report will show that there is one area of "high" fracture intensity, and two areas of "low" fracture intensity (green). The general interpretation is that the low fracture intensity areas are underlain by topographic highs on the Basement. With this established, the deduction is that the Basement is high in the central-east and southwest parts of Permit 5059.

These Basement high features are most interesting from the oil and gas point of view. The general shape of the features is such that the causative feature must be a hill on the Basement surface. A fault is unlikely as the causative feature as the high areas are greater than one and one-half miles in width. If a fault caused the fracture "low" the width of the low would be less than one mile.

Three hypothetical structure cross-sections accompany this report and reference to them will show how Basement "highs" are inferred to be present beneath areas of low fracture intensity. Two profiles run at right angles to the strike of the Basement while the third is parallel to strike.

Respectfully submitted by:

RAYALTA PETROLEUMS LTD.

William H. Crook

WGC/lp

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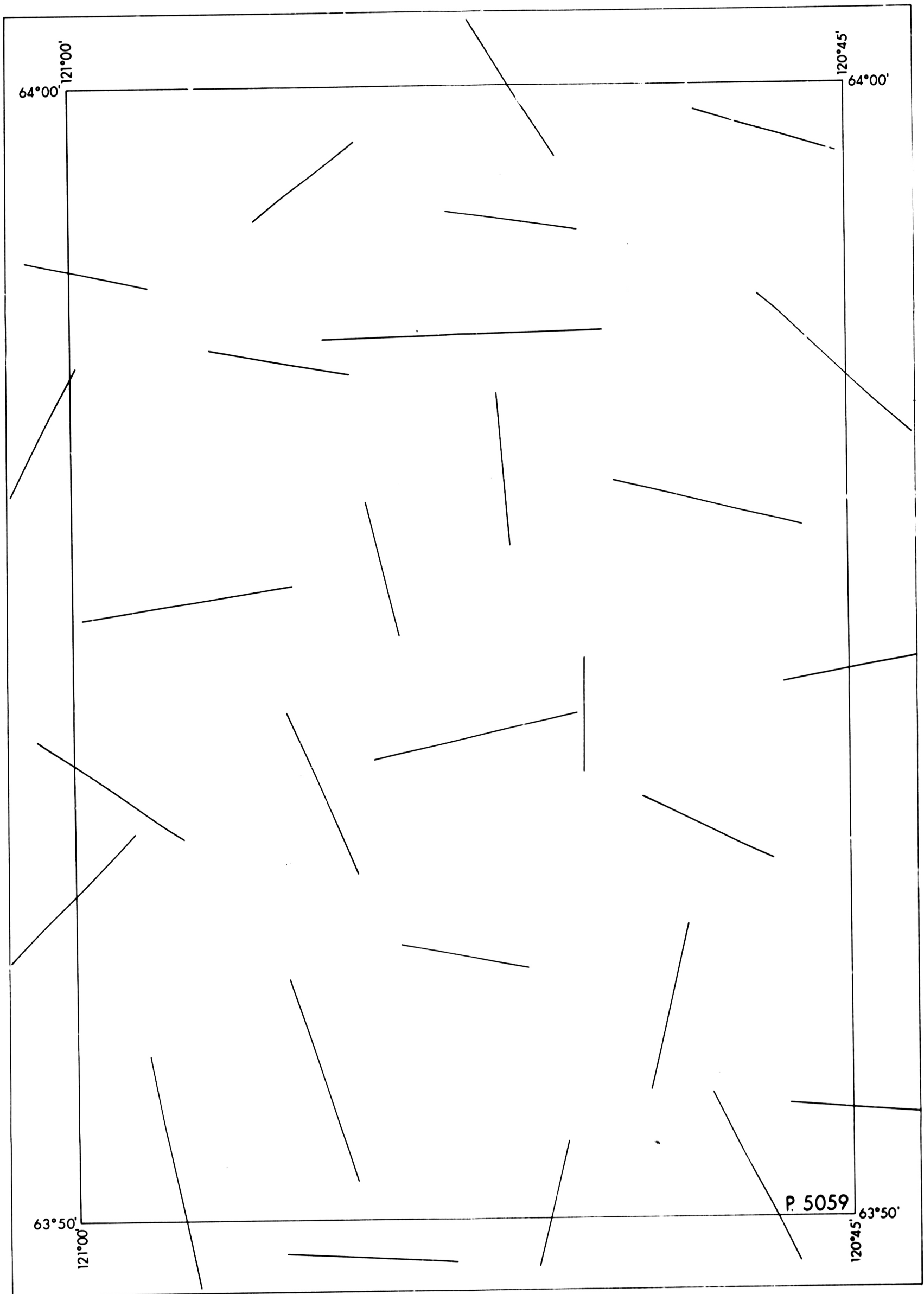
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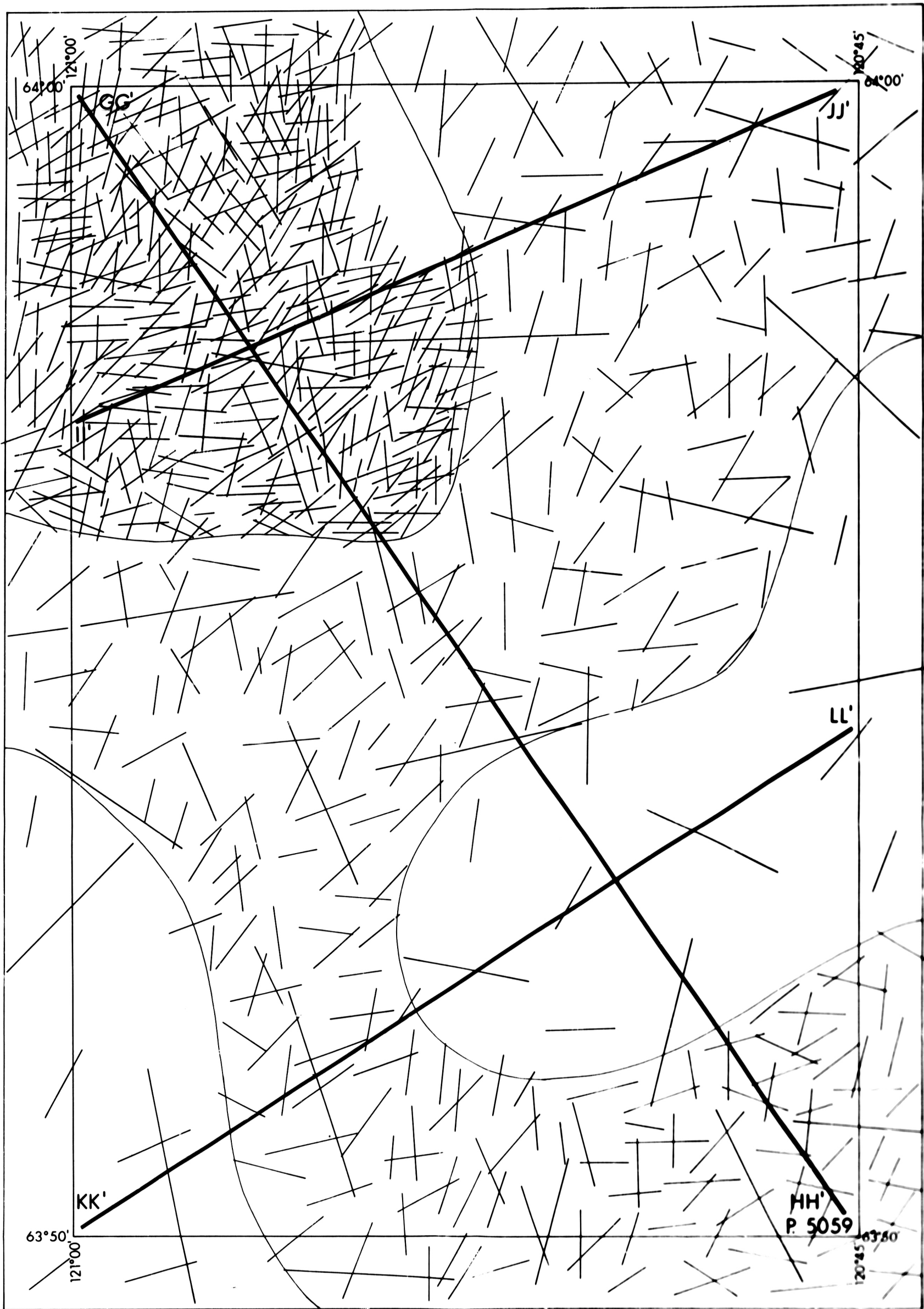
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MEGA FRACTURE PATTERN 662-1-5-12

SCALE IN MILES





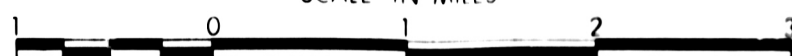
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TOTAL FRACTURE PATTERN

662-1-8-12

SCALE IN MILES



- LOW DENSITY
- NORMAL DENSITY
- HIGH DENSITY

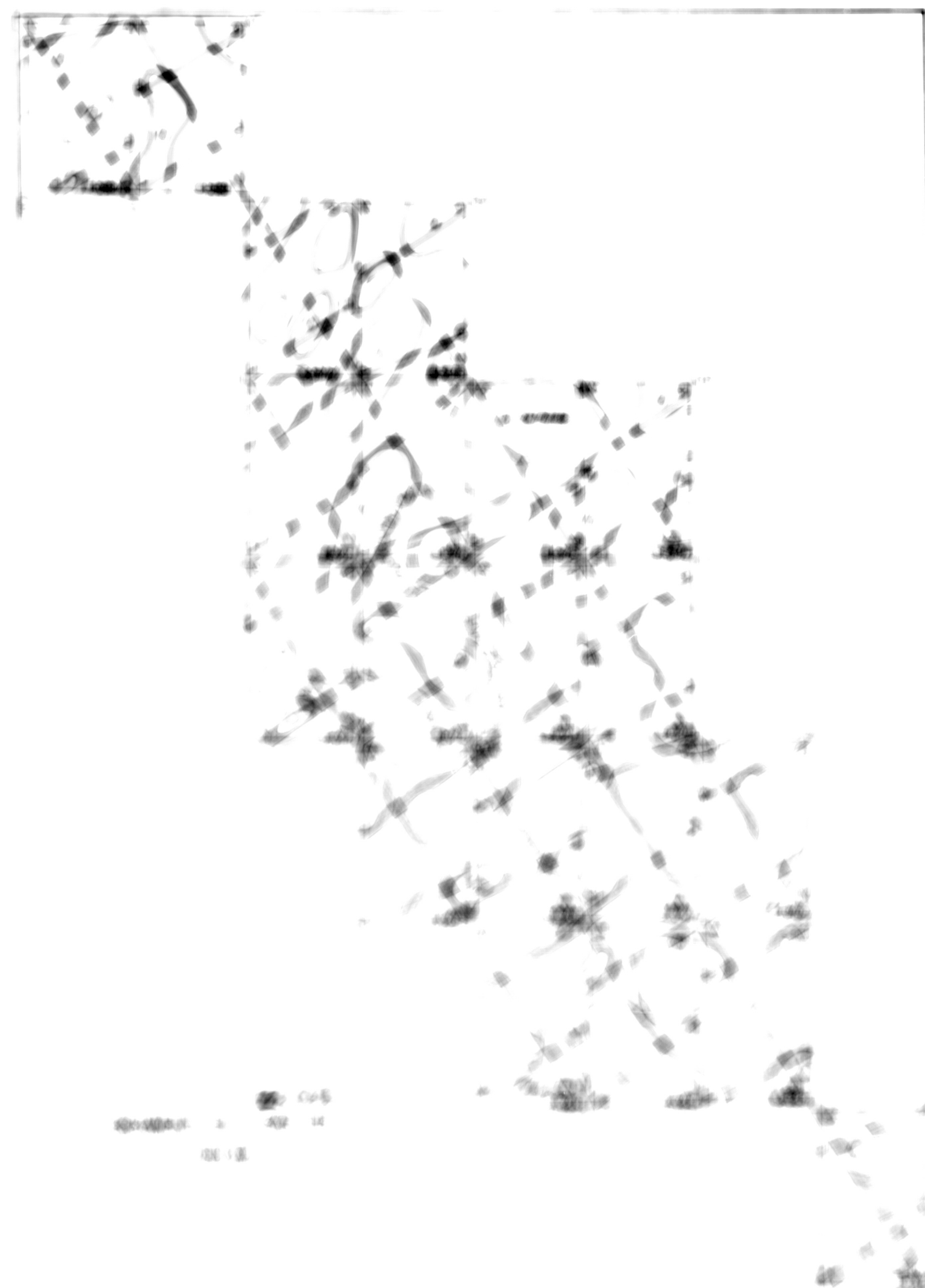


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比例 1:1
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GENERAL GEOLOGY
to
FRACTURE ANALYSIS SURVEY

of
P N G PERMIT NO. 3039

for
GROSMONT OIL & GAS LTD.

by
RAYALTA PETROLEUMS LTD.

INTRODUCTION

This report discusses the results of a Fracture Analysis Survey carried out within, and in the immediate vicinity of, Petroleum and Natural Gas Permit No. 5059. This Permit is located in the Northwest Territories and is held under the Canada Oil and Gas Land Regulations and is located between $120^{\circ} 45'$ to $122^{\circ} 00'$ longitude and $63^{\circ} 50'$ to $64^{\circ} 00'$ latitude. The Permit is 760 miles north of Edmonton and 240 miles northwest of Yellowknife.

The Yellowknife Highway serves Fort Providence which is 200 miles southeast of the Permit and is the closest road to the area. Access to the Permit itself is by helicopter or on foot during the summer or by vehicle during the months when the ground is frozen. However, there are no roads in the area and considerable road construction would be required to reach any particular area. Some narrow cut lines are present and these afford limited access.

The surface of the Permit is quite flat-lying and total relief does not exceed 100 feet. There is only a poorly developed drainage pattern within this area and a few intermittent streams flow northwest towards the Johnny Hoe River. A layer of very soft muskeg covers this part of the Northwest Territories and this muskeg is so soft that it is impassable to all but specialized vehicles.

Vegetation consists of thick stands of thin evergreen trees interspersed with many small open areas. These open areas are covered by muskeg grass and scrub deciduous growth. The evergreen trees show up as a medium gray tone on the mosaic and the open areas are a lighter gray. A few small patches of deciduous trees are present

There is no topographic form or aerial photo feature present which immediately suggests the presence of any geologic structure.

The results of this survey are illustrated
on the Total Fracture Map, the Stage Fracture
Map plus the maps with the fractures super-
imposed. In addition there are three suggested
cross sections. All the above can be found in
the folder at the back of this report.

STRATIGRAPHY

GENERAL STATEMENT

The stratigraphic discussion presented herewith is based on a study of the area covered by Petroleum and Natural Gas Permit 1053 and 1054 to 1062 inclusive. The north limit of this area is located along the south shore of the North Arm of Great Bear Lake and extends eastward to about 64° 00' - 127° 00'. No wells have been drilled in this area and surface outcrops are rare and widely scattered. Therefore, it has been necessary to study the regional geology of the whole area through reconnaissance and make many projections of data and, admittedly, some of these projections are rather long ranged. However, when combined with such subsurface information as is available an approximate picture of the sedimentary stratigraphy can be presented.

The Permits are on the Interior Plains 40 to 60 miles east of the Franklin Mountains, and about 100 miles west of the Pre-Cambrian Shield outcrop area. The Permits lie about 150 miles east-southeast of the Norman Wells Oil Field which provides most of the nearest well control. The area covered by the above referred Permits is underlain by sediments ranging in age from Cambrian to Tertiary. Regional isopachs indicate about 6,000 feet of sediments should be present under the northern Permits and about 4,500 feet under the southern Permits. Structurally, they should be underlain by homoclinal to gently folded beds. However, since the Northern Franklin Mountains are the result of compressional movements, it would seem likely that the area lying in front of the mountains may have been folded into anticlines of appreciable magnitude. The Pre-Cambrian may have undergone early faulting under the Permits since faulting of this age is quite common in this region. The Basement

faults have commonly produced northeast trending lineaments but are generally conceded to have not disturbed the Paleozoic sediments. The Basement faults are generally steeply inclined right-hand faults and as far as is known, the horizontal movement exceeds the vertical movement by a large amount. Some recurrent movement at widely separated times has been noted in the region. Well control east of the acreage concerned is very scarce. Regional isopachs and facies maps along with published geological reports have been used to describe the stratigraphic sequence which might be expected to underlie this area.

CAMBFIAN and/or OLDER

KATHEFINE GFOUP

The Katherine Group which represents the earliest Paleozoic sediments in this region, is named from a section exposed in the Upper Carcajou River area which lies about 150 miles

west of the acreage under study. The section exposed consists of interbedded quartzites and black, platy shales. The shales which are black, platy, bituminous as well as green and chocolate coloured, are contained in interbeds within the quartzites. The quartzites are generally pink, buff, rusty and white in outcrop. The top of the Katherine is placed at the base of a chocolate coloured shale succession while the base was not seen in outcrop leaving the total thickness unknown for this area. The Katherine Group has not been penetrated by any drill holes in this region to date, which means the subsurface section is unknown. While reservoir beds are not described in outcrop it must be expected that sand bodies such as offshore bars, beach sands and long shore bars will eventually be found in this group of sediments. Similar sands are found to be prolific producers in the Red Earth Creek area of northern Alberta. The delineation of productive areas for and underlying such sands is dependant

on a knowledge of present Pre-Cambrian structure as well as its topographical expression, when the sands were being deposited. A gravity meter and airborne magnetometer survey could be used to good advantage in locating areas for more detailed exploration. Source rocks for hydrocarbons should be no problem since the outcrop section previously described would appear to contain an adequate source within its bituminous shales. This section should be considered in any exploratory plans for this area.

CAMBRIAN

MACDOUGAL GROUP

The type section of the Macdougall Group is located about 130 miles west of this area in the Dodo Canyon of the Macdougall River. At the type section the Macdougall is divisible into a number of formations which total 997 feet in thickness. The base is placed at the bottom

of a 130 foot thick chocolate brown shale while the top is placed above 50 feet of evenly bedded limestone with shale partings. The lithology is made up of interbedded limestones, sandstones, reddish coloured gypsum, black, petroliferous shales, red and green shales as well as chocolate coloured shales. The Imperial River section which was mapped by Laudon lies 30 miles to the northwest of the type section. The section, which is 1,839 feet thick with the base not exposed, consists of alternating sandstones, limestones, gypsum and vari-coloured shales. The lower part consists of sandstones with minor shale interbeds which appear to be a shallow water deposit since they are ripple marked and cross-bedded. The section becomes increasingly shaly upwards. The gypsum content is also greater near the top. A 146 foot thick bed of black to dark grey, laminated, algal limestone is located near the top of the section. Calcareous algae up to three feet in diameter are present. At

Norman Wells the Macdougall Group contains a bed of salt 2,000 feet thick which is correlated with the Saline River Formation. This salt section is believed to be present to the north, west and south of Norman Wells for the following reasons:

- 1 The Western margin of the Saline River salt is known in the Norman Wells area and a postulated extension of this margin can be made to the north, west and south of Norman Wells.
- 2 The overlying Ronning carbonates are brecciated at exposures in the northern Richardson Mountains west of Inuvik, suggesting salt solution collapse.
- 3 The type section at Saline River which lies 100 miles south of

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creates the possibility of salt structures in the overlying carbonate banks similar to those found to be productive in southeast Saskatchewan and at Rainbow Lake in northwestern Alberta. The algal laminate at Imperial River indicates some organic activity in the Macdougall seas and this coupled with underlying salt features, could give rise to hydrocarbon bearing reservoirs within this sequence. The petroliferous shales within the Macdougall should be adequate source material. The Macdougall has been reached by very few of the wells drilled in this region and no where has it been fully penetrated. Imperial Vermilion Ridge No. 1, drilled 3,177 feet of Macdougall beds without reaching the underlying Katherine Group. To date no reservoirs have been tested in the wells which have drilled to the Macdougall.

ORDOVICIAN-SILURIAN

RONNING FORMATION

Rocks of Ordovician Age have not, as noted by various authors, been definitely identified in this region; however, it seems to be generally accepted that they are present in the Norman Wells region. The contact with the underlying Macdougall is unconformable. Stetck mapped 1,500 feet of shales and argillites at outcrops in the Upper Peel River area, which lies some 300 miles to the west of these Permits. About 150 miles west of the Permits, at the Keele and Twitya River confluence, the Ordovician section was mapped by Keele as 4,000 feet of alternating beds of argillite, dolomite and limestone with 1,500 feet of sandstone overlying and separated from them by a 100 foot thick diabase sill. He mapped this same sandstone 35 miles to the east as being 4,500 feet thick with only occasional shale partings. The sections described in outcrop by Keele

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Figure 1. The effect of the concentration of the inhibitor on the rate of polymerization of α -methylstyrene in the presence of SnCl_4 at 25°C .

shells and corals composed of porosity

A. The nearest outcrop section of the
formation which is located about 50 miles west
of the town of St. Charles on Great
Bear River the Franklin Mountain Formation
is about 885 feet in thickness. The section
consists of limestone with the basal 200 feet
described as carbonaceous about midway in the
section is 25 feet of cherty limestone. The upper
25 feet is a grey, dolomitic limestone. The
base of the Franklin Formation here is not
exactly clear as various workers have included
beds beneath those described in the Ranney
(2) and so on. They consist of gypsum,
carbonaceous limestone with black, bituminous
shales and highly bituminous limestones, which
some have had the thought to the writer. The
thickest of units consists of 400 feet of carbon-
aceous the upper 25 feet is a dolomitic lime-
stone. The remaining 200 feet which overlying it are
85 feet of limestone and other beds. The upper

180 feet is a hard, grey dolomite that is cherty in the lower part. The section is overlain by the Bear Rock brecciated dolomites. The section at Bear Rock near Fort Norman, which is 30 miles west of Mt. St. Charles, consists of 600 feet of limestone, dolomites and shales with the brecciated sediments of the Bear Rock overlying them and the Macdougall red and green, gypsiferous shales underlying them. The Mount Kindle is apparently not present here. Imperial Loon Creek No. 2, in $65^{\circ} 07' 20''$ N., and $126^{\circ} 12' 51''$ W., which is about 75 miles west of the Permits, penetrated 1,270 feet of Ronning which is close to the same thickness as mapped at Mt. St. Charles. The Loon Creek well found the Ronning to consist mainly of white to grey, micro-crystalline to granular dolomites with some evaporitic plugging. Scattered poor porosity was present throughout; however, no tests were run. Outcrops of the Ronning are found about 200 miles to the northwest of the Permits along the Hare Indian

River. The section consists of 750 feet of limestones, overlain by the Bear Rock with the base not exposed. The section is not identified as Mount Kindle but regionally it should be present at this location.

Stelck mapped 100 feet of massive, crystalline, porous limestones containing some coralline fauna at Schooner Creek, which is four miles north of Norman Wells. He correlated them with the lower portion of the Mount Kindle Formation. This section can be interpreted as a porous, carbonate bank deposit. The Mount Kindle is likely to have a number of these carbonate banks or low transgressive reef fronts in this area, since, as can be seen from the various sections described above, it undergoes both facies changes and thickness changes in this region. Since the Mount Kindle is present on Mt. St. Charles to the east of the Permits, as well as to the north of them it will doubtlessly be present under

(b) As outlined previously, low reef fronts or carbonate banks may be present and coupled with a seal provided by overlying Bear Rock evaporites could present an extensive trap. Lateral facies changes from porous to semi-evaporitic carbonates also provide a potential trap of considerable areal extent.

(c) Selective solution of the underlying Cambrian Saline River salt may give rise to one or two stage solution structures such as are found to be productive of oil in the Mississippi River area of south-east Queensland. Partial solution of the salt prior to or during stage oil production would have resulted in the formation of solution structures in the salt beneath which the oil was trapped. It is suggested that these structures are

would provide the loci for reef
and/or carbonate banks to grow
on. Traps of the Hummingbird
type would involve early local
solution of the salt. This may
have occurred in late Cambrian
or early Ronning time. The
depressions created would receive
an extra fill of sediments over
that being deposited where the
salt was not removed. Once
sedimentation within the sink
caught up, subsequent sediments
would be deposited on a normal
sea floor. The second stage in
the formation of the Hummingbird
type trap would involve the re-
moval of the salt surrounding the
original sink at some time after
that no longer results in solution.
This would leave the trap in place
with the surrounding sea floor of the

original salt solution structurally high. The Bear Rock evaporites should provide an effective reservoir seal. Evidence to support one or two stage salt removal in this region is present in the brecciated nature of the sediments composing the Lower Ronning and Bear Rock sediments in known sections.

(d) Gentle to tight anticlinal folds may have been formed by some of the numerous periods of structural activity which have occurred in this region.

MIDDLE DEVONIAN

BEAR ROCK FORMATION

The Bear Rock Formation overlies the Ronning Group and is separated from it by a

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Figure 1. The effect of the concentration of the *Agrobacterium* suspension on the transformation efficiency of *Agrobacterium* strains.

Figure 1

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姓名: 王 强 性别: 男 年龄: 25 岁 职业: 教师 单位: 某某中学 电话: 13800138000 电子邮箱: wangqiang123@163.com

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Mountains continues southeast along the western side of the MacKenzie Mountains. The basinal shales are flanked by a belt of shelf-edge limestones and dolomites along their eastern side. Porosity is developed within these carbonates. Adjacent to the shelf-edge carbonates and covering much of the Interior Plains and Peel Plateau area are the shelf limestone and dolomite facies. In the Peel Plateau they attain a thickness of some 2,000 feet and consist of micritic, pellet and micritic skeletal limestone with intervals of finely crystalline, porous dolomite in the lower part. The shelf carbonates are in turn replaced by a relatively narrow belt of shelf dolomites. This takes place in the MacKenzie Mountains and extends in a line north through the Fort Good Hope region and south into the Camell and Nahanni Ranges. The shelf dolomites in turn are replaced by an evaporite facies along their entire length. This facies change begins to the west of Norman Wells. In the Norman

Wells area and also in the area of the Permits under discussion the basal portion of the Bear Rock is commonly evaporitic while the upper portion consists of carbonate breccias. The evaporite facies extends southward into northern Alberta where it is known as the Chinchaga Formation. South of Norman Wells a strong depositional feature called the Camsell Basin occurs. Thickening from 2,000 feet to more than 5,000 feet, accompanied by facies changes from evaporites through shelf carbonates to basinal sediments takes place into this basin. The shelf carbonates are cryptocrystalline to microcrystalline dolomites while the shelf-edge facies is reefal with some of it at least being porous.

The Bear Rock carbonates in the Norman Wells area have been found to be very porous in some wells while in others the porosity has been plugged by anhydrite and gypsum. Considerable bitumin has been en-

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ever, this does not seem very plausible. The rapid facies changes which may be expected within the Bear Rock is evident when the Mt. St. Charles section is compared to an exposure three miles further north. Here the chert beds of the underlying Mount Kindle Formation are overlain by 500 feet of grey gypsum beds that are in turn overlain by limestone beds that are mapped as part of the overlying Hume (Ramparts) Formation. The Mt. St. Charles section has no evaporites. The thickness variation between these two sections is worthy of comparison also.

HUME FORMATION

Considerable confusion has existed in the literature concerning the relationship of the Ramparts or Hume, Hare Indian and Kee Scarp Reef. A paper by H.G. Basset in the Geology of the Arctic Symposium is probably the most important to an understanding of the Middle Devonian geology of this area.

Hume defined the Ramparts Formation as containing all definite Middle Devonian beds in the Norman Wells area as well as in the surrounding area. The base would be placed at the top of the underlying Bear Rock and the top at the contact with the overlying Fort Creek Shales. He divided the Ramparts into three members, a lower limestone member, a middle shale member and an upper limestone member. The lower limestone, which is relatively thin in the Norman Wells area, thickens in a northwest direction. About 60 miles west-northwest of Norman Wells in the Imperial Range on Mountain River the Lower Ramparts is described as 445 feet of limestone, dark grey to black, with irregular black shale partings, very fossiliferous in part (particularly corals) and very petroliferous in part. The Middle Ramparts of this section consists of 700 feet of grey to green shales and limey shales with many thin limestone beds which are commonly conquinoid in the lower

കേരളത്തിലെ മറ്റൊരു പ്രവിശ്യയായ തൃശ്ശൂരിൽ ൧൯൪൭-ൽ

ഒരു കുടുംബത്തിൽ ജനിച്ചു. അദ്ദേഹത്തിന്റെ പേര് കെ. ജി. ജോർജ്ജ്.

തൃശ്ശൂരിൽ തന്നെ തന്റെ പ്രാഥമിക വിദ്യാഭ്യാസം പൂർത്തിയാക്കി.

൧൯൫൫-ൽ തൃശ്ശൂർ കോളേജിൽ നിന്ന് ബിരുദം നേടി.

അദ്ദേഹത്തിന്റെ തുടർ വിദ്യാഭ്യാസം മദ്രാസ് കോളേജിൽ തുടർന്നു.

മദ്രാസ് കോളേജിൽ നിന്ന് ബിരുദം നേടി. തുടർന്ന് ൧൯൫൭-ൽ

കേരളത്തിൽ തന്നെ തന്റെ വിദ്യാഭ്യാസം തുടർന്നു. ൧൯൫൯-ൽ

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productive Keg River pinnacle reefs in north-western Alberta. The Keg River platform is almost certainly the source of the oil in these prolific reefs, and because of the similarities outlined above any reservoirs developed in the Hume must be considered as prospective.

HARE INDIAN

The contact of the Hare Indian with the underlying Hume is generally sharp and probably represents a sudden influx of mud into a clean well aerated sea. It appears to represent a mud bank deposit with the source area lying to the northeast, partially filling a large basin. The contact of the Hare Indian with the overlying Kee Scarp is somewhat diachronous, since it is generally placed at the point the section changes from predominant shale to predominant limestone. Facies changes thus account for the diachronous nature of the contact as well as having been the cause of some of the confusion which has

Figure 10-18

Figure 1. The effect of the concentration of the *Agrobacterium* strain on the transformation efficiency of *Agrobacterium* strain.

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● 國 家 領 土 主 權 的 確 立 ●

174명 / 1월 7일 / 1월 11일 / 1월 15일 / 1월 19일 / 1월 23일 / 1월 27일 / 1월 31일

● 附 录 ●

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1948年 1月 2日 星期三 晴 1月 2日 星期三 晴

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Figure 6. The effect of the initial concentration of the monomer on the polymerization rate.

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Journal of Management Education 36(8) 907-924

● 2010 年 12 月 1 日起, 凡在 2010 年 12 月 31 日前, 已持有《二手车交易统一发票》, 且该发票所载日期在 2009 年 1 月 1 日至 2009 年 12 月 31 日期间的, 其二手车交易环节应缴纳的二手车交易税, 暂免征收。

FREE SCARF

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Scarp in the Norman Wells area consists of a lower platform unit which is about 75 feet to 165 feet thick and lithologically is a bedded limestone with abundant fossils. The platform unit is usually devoid of hydrocarbons. Overlying the foundation unit is a biohermal reef which constitutes the reservoir for the Norman Wells Oil Field. The reef is composed of materials such as corallites, bryozoans and stromatopoids set in a coral sand matrix. The facies varies widely between wells as would be expected in a true reef. The thickness of the Kee Scarp reef above the platform unit varies from zero (0) feet to 350 feet in the Norman Wells area. The greatest overall measured thickness of Kee Scarp in the area is 495 feet. The Kee Scarp is overlain by the Canol Formation, or, in its absence, the Fort Creek shales which Basset redefined as part of the Imperial Formation.

The oil in the Norman Wells Field is trapped in the upper end of a discrete Kee Scarp reef. The thickness of the reef ranges up to a total of 495 feet. Reserves in the reef have been estimated as high as 60,000,000 barrels while the productive area of the field is placed at 2,600 acres.

The platform unit of the Kee Scarp is undoubtedly the correlative of the Upper Farnorts limestone unit mapped by Hume. This fact, as mentioned above, means the Kee Scarp is a widespread unit. Since the Kee Scarp reef grows upwards from the platform unit any well drilled in this area and any acreage held, must be considered as possibly containing discrete Kee Scarp reefs. Maximum reef growth, regionally, has generally been found on the margins of Hare Indian thickets, however, the presence of them does not ensure Kee Scarp reefs. The margins of the two Hare Indian thickets, which were described

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附錄一

● 2010年10月1日

1. 1990 1991 1992 1993 1994 1995 1996 1997 1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013 2014 2015 2016 2017 2018 2019 2020 2021 2022 2023 2024 2025 2026 2027 2028 2029 2030 2031 2032 2033 2034 2035 2036 2037 2038 2039 2040 2041 2042 2043 2044 2045 2046 2047 2048 2049 2050 2051 2052 2053 2054 2055 2056 2057 2058 2059 2060 2061 2062 2063 2064 2065 2066 2067 2068 2069 2070 2071 2072 2073 2074 2075 2076 2077 2078 2079 2080 2081 2082 2083 2084 2085 2086 2087 2088 2089 2090 2091 2092 2093 2094 2095 2096 2097 2098 2099 2100 2101 2102 2103 2104 2105 2106 2107 2108 2109 2110 2111 2112 2113 2114 2115 2116 2117 2118 2119 2120 2121 2122 2123 2124 2125 2126 2127 2128 2129 2130 2131 2132 2133 2134 2135 2136 2137 2138 2139 2140 2141 2142 2143 2144 2145 2146 2147 2148 2149 2150 2151 2152 2153 2154 2155 2156 2157 2158 2159 2160 2161 2162 2163 2164 2165 2166 2167 2168 2169 2170 2171 2172 2173 2174 2175 2176 2177 2178 2179 2180 2181 2182 2183 2184 2185 2186 2187 2188 2189 2190 2191 2192 2193 2194 2195 2196 2197 2198 2199 2200 2201 2202 2203 2204 2205 2206 2207 2208 2209 2210 2211 2212 2213 2214 2215 2216 2217 2218 2219 2220 2221 2222 2223 2224 2225 2226 2227 2228 2229 2230 2231 2232 2233 2234 2235 2236 2237 2238 2239 2240 2241 2242 2243 2244 2245 2246 2247 2248 2249 2250 2251 2252 2253 2254 2255 2256 2257 2258 2259 2260 2261 2262 2263 2264 2265 2266 2267 2268 2269 2270 2271 2272 2273 2274 2275 2276 2277 2278 2279 2280 2281 2282 2283 2284 2285 2286 2287 2288 2289 2290 2291 2292 2293 2294 2295 2296 2297 2298 2299 2300 2301 2302 2303 2304 2305 2306 2307 2308 2309 2310 2311 2312 2313 2314 2315 2316 2317 2318 2319 2320 2321 2322 2323 2324 2325 2326 2327 2328 2329 2330 2331 2332 2333 2334 2335 2336 2337 2338 2339 2340 2341 2342 2343 2344 2345 2346 2347 2348 2349 2350 2351 2352 2353 2354 2355 2356 2357 2358 2359 2360 2361 2362 2363 2364 2365 2366 2367 2368 2369 2370 2371 2372 2373 2374 2375 2376 2377 2378 2379 2380 2381 2382 2383 2384 2385 2386 2387 2388 2389 2390 2391 2392 2393 2394 2395 2396 2397 2398 2399 2400 2401 2402 2403 2404 2405 2406 2407 2408 2409 2410 2411 2412 2413 2414 2415 2416 2417 2418 2419 2420 2421 2422 2423 2424 2425 2426 2427 2428 2429 2430 2431 2432 2433 2434 2435 2436 2437 2438 2439 2440 2441 2442 2443 2444 2445 2446 2447 2448 2449 2450 2451 2452 2453 2454 2455 2456 2457 2458 2459 2460 2461 2462 2463 2464 2465 2466 2467 2468 2469 2470 2471 2472 2473 2474 2475 2476 2477 2478 2479 2480 2481 2482 2483 2484 2485 2486 2487 2488 2489 2490 2491 2492 2493 2494 2495 2496 2497 2498 2499 2500 2501 2502 2503 2504 2505 2506 2507 2508 2509 2510 2511 2512 2513 2514 2515 2516 2517 2518 2519 2520 2521 2522 2523 2524 2525 2526 2527 2528 2529 2530 2531 2532 2533 2534 2535 2536 2537 2538 2539 2540 2541 2542 2543 2544 2545 2546 2547 2548 2549 2550 2551 2552 2553 2554 2555 2556 2557 2558 2559 2560 2561 2562 2563 2564 2565 2566 2567 2568 2569 2570 2571 2572 2573 2574 2575 2576 2577 2578 2579 2580 2581 2582 2583 2584 2585 2586 2587 2588 2589 2590 2591 2592 2593 2594 2595 2596 2597 2598 2599 2600 2601 2602 2603 2604 2605 2606 2607 2608 2609 2610 2611 2612 2613 2614 2615 2616 2617 2618 2619 2620 2621 2622 2623 2624 2625 2626 2627 2628 2629 2630 2631 2632 2633 2634 2635 2636 2637 2638 2639 2640 2641 2642 2643 2644 2645 2646 2647 2648 2649 2650 2651 2652 2653 2654 2655 2656 2657 2658 2659 2660 2661 2662 2663 2664 2665 2666 2667 2668 2669 2670 2671 2672 2673 2674 2675 2676 2677 2678 2679 2680 2681 2682 2683 2684 2685 2686 2687 2688 2689 2690 2691 2692 2693 2694 2695 2696 2697 2698 2699 2700 2701 2702 2703 2704 2705 2706 2707 2708 2709 2710 2711 2712 2713 2714 2715 2716 2717 2718 2719 2720 2721 2722 2723 2724 2725 2726 2727 2728 2729 2730 2731 2732 2733 2734 2735 2736 2737 2738 2739 2740 2741 2742 2743 2744 2745 2746 2747 2748 2749 2750 2751 2752 2753 2754 2755 2756 2757 2758 2759 2760 2761 2762 2763 2764 2765 2766 2767 2768 2769 2770 2771 2772 2773 2774 2775 2776 2777 2778 2779 2780 2781 2782 2783 2784 2785 2786 2787 2788 2789 2790 2791 2792 2793 2794 2795 2796 2797 2798 2799 2800 2801 2802 2803 2804 2805 2806 280

1. 關於本會之組織及職權範圍

第一章 總則

第一條 本會定名為「中華民國醫學會」

第二條 本會之宗旨在促進醫學學術之進步及醫藥衛生事業之發展

第三條 本會之組織由會員大會及常務理事會組成

第四條 本會之會員分為正式會員、名譽會員及通訊會員三種

第五條 凡具有中華民國國籍之醫學博士或具有同等學力者

第六條 凡具有中華民國國籍之醫學士或具有同等學力者

第七條 凡具有中華民國國籍之醫學士或具有同等學力者

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第十七條

CRETACEOUS

SANS SAULT GROUP

The Sans Sault Group is the basal group of Cretaceous sediments which lie directly above the disconformity separating Cretaceous and Devonian sediments. The top of the Group is usually placed at the base of the first bentonite bed in the overlying thick shale sequence. The sequence consists of shales and sandstones of marine origin. The thickness is about 1,411 feet at the Sans Sault section.

SLATER RIVER FORMATION

The Slater River which overlies the Sans Sault Group, consists of thin bedded, silty, friable shales with abundant ironstone concretions. There are also some beds of white and yellow alum and sulphur. Sandstone is only occasionally present. There are many beds of bentonite, which in outcrops are often visible. The Slater River

Formation also contains a fish scale horizon which is thought to indicate an upper Cretaceous Age for the formation. This formation is about 1,000 feet thick at the type section.

LITTLE BEAR FORMATION

The type section of this formation is west of Fort Norman on the Little Bear River. The beds consist of sandstone, some conglomerates, sandy shales and coal seams. The beds are not correlatable between areas due to their lenticular nature. The beds are 780 feet thick at their type section and contain marine, brackish and fresh water fossils.

EAST FORK FORMATION

The East Fork Formation is made up of a series of well stratified, grey, conchoidal and plastic marine shales. There are some thin limy sandstone members and thin coal seams near the base. The thickness of this formation is 850 feet at its type locality on the

East Fork of the Little Bear River.

The thickness of Cretaceous beds present underlying this Permit area is very difficult to ascertain. C.S. Lord noted coal deposits on Etacho Point which lies about 100 miles due north of the Permits on the west side of Great Bear Lake. The coal which is lignite, is contained in about 1-1/2 miles of outcrops. The outcrops usually contain several seams separated by a few feet of clay, sand, or silt. The width of one seam is from 12 feet to 17 1/2 feet and may be about 7,000 feet in length. The age of the coal is not given, but it may be part of the Little Bear Formation.

The unconformity which underlies the Cretaceous in this area has probably removed much of the Devonian Imperial Formation from the area covered by these Permits. Since subsurface and surface control is so sparse in this area, any prediction of the depth of

this erosion is very difficult to make. North of Norman Wells this erosion has in places removed the entire Upper Devonian sequence, leaving the Middle Devonian Formation at subcrop.

TERTIARY

The Tertiary sediments in the Norman Wells area are not subdivided. They consist of conglomerates, gravels, shales, lignites, soft, coarse, carbonaceous sands and soft clays. The Tertiary is exposed south of the Permits under review in the Mt. St. Charles area along the Great Bear River. Plants collected from the exposures along the Great Bear River indicate an Eocene Age. The thickness is approximately 600 feet at these exposures. At exposures on the Little Bear River, 1,600 feet of Tertiary sediments have been mapped. Near the headwaters of the East Fork River beds up to 1,200 feet

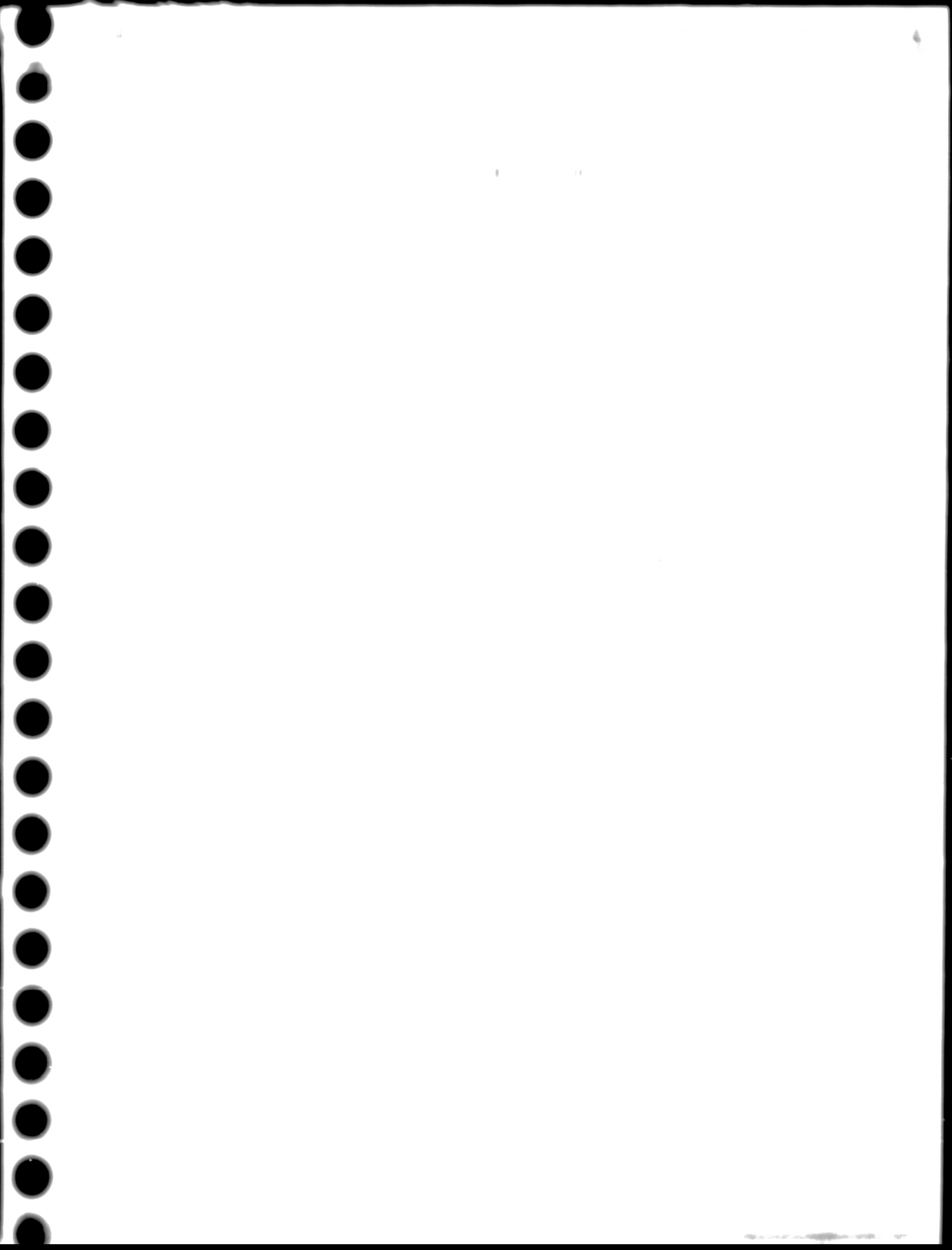
have been mapped with coal seams eight feet to ten feet thick. The sections mentioned form part of a basin which dips to the southwest in this area.

It is recommended that further evaluation of the Permits under review consist of gravity meter and/or airborne magnetometer surveys. They should be of great assistance in outlining the distribution of the Saline River salt and any salt structures associated with it. The present structure of the Pre-Cambrian Basement could probably be mapped by this method, also, as well as providing a better idea of the drilling depth to it.

FRACTURE ANALYSIS

This section of the report discusses the results of a Detailed Fracture Analysis Survey carried out on the area under discussion. An aerial mosaic (scale 1.5 inches equals approximately 1 mile) made from Dominion Government aerial photographs accompanies this report. These same photographs were examined stereoscopically and the fractures plotted on the individual photographs, then transferred to the mosaic for analysis.

The theory that the earth's crust is abundantly and methodically fractured is the basic premise on which is built the exploration technique known as Fracture Analysis. A Fracture is defined as "... generally abundant, natural lineation discernible on aerial photographs"



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解: 因为 $\frac{1}{2} \times 100 = 50$, 所以 $\frac{1}{2}$ 吨 = 50 吨。

Figure 1. The effect of the concentration of the inhibitor on the rate of polymerization of α -methylstyrene in the presence of SnCl_4 at 50°C .

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INTERPRETATION OF FRACTURE DATA

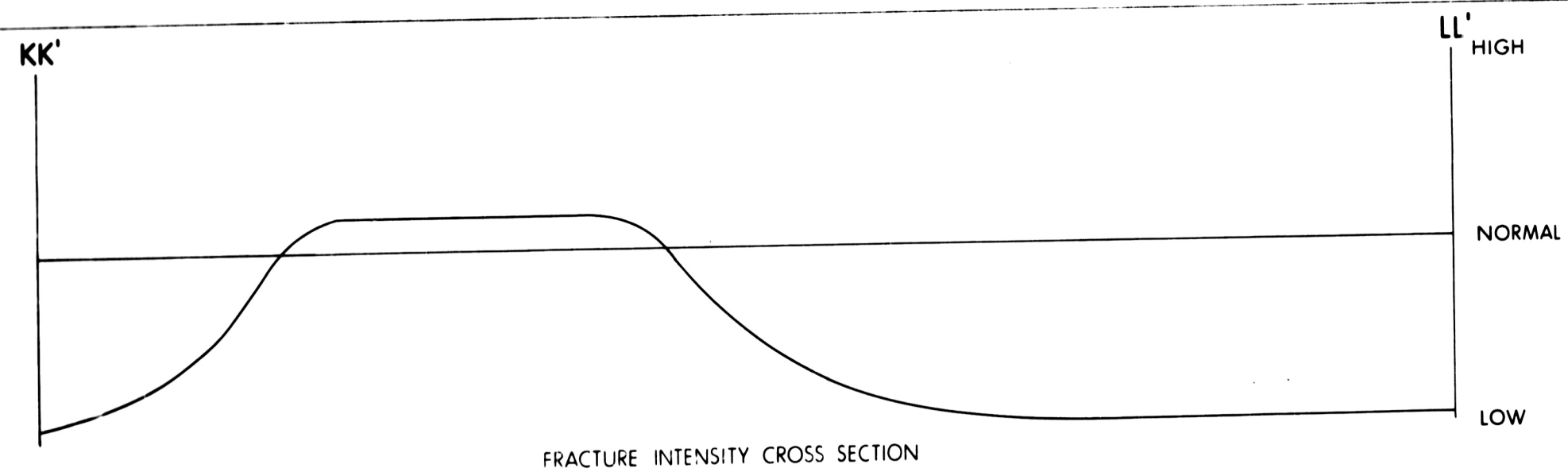
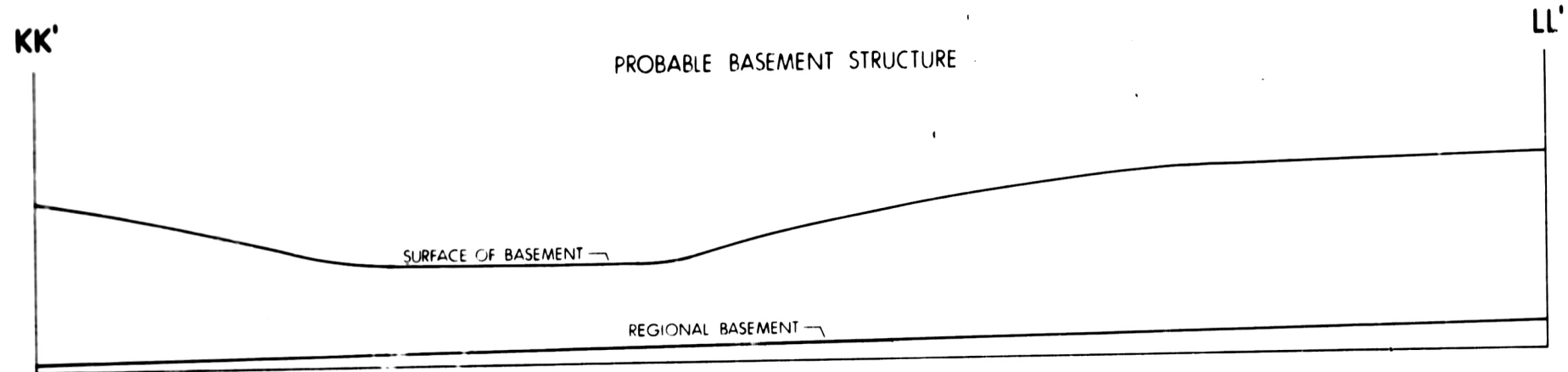
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Because of certain inherent limiting factors, Structure Incidence Surveys have a lower order of reliability than Detailed Fracture Analysis Surveys. To some extent at least, surface conditions affect the fracture count. In areas covered by lakes, sloughs and rivers, the fracture count is zero. Cultivated areas generally yield a lower count than adjacent virgin territory. Consequently, a difference or contrast in fracture counts (F/I) between two points may be in part due to surface type. But even this is not so different surface conditions. For example, the same surface may be covered by different types of vegetation, or the same type of vegetation may be found on different surface types. These factors may also affect the fracture count.

Therefore, the fracture count is not a true measure of the amount of fracture in a given area. It is only a relative measure. The fracture count is a measure of the relative amount of fracture in a given area compared to the amount of fracture in a given area. The fracture count is a measure of the relative amount of fracture in a given area compared to the amount of fracture in a given area. The fracture count is a measure of the relative amount of fracture in a given area compared to the amount of fracture in a given area.

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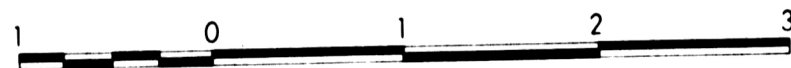


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